INUNDATION AND EROSION SUSCEPTIBILITY IN WAVE DOMINATED BEACHES

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Abstract – Hydrodynamic forces over the beach sediments are the main driving factors affecting the frequency and magnitude of morphological changes in beach systems. Most of the time, these driving factors act in a foreseeable way and do not represent any danger either to the coastal systems or to their population. However, hydrodynamic forces are also capable of inducing high morphodynamic behaviour on the beach profiles and very often in a short period of time; this endangers people and property and leads to system retreat. The most common consequences of the occurrence of this type of phenomena over the landforms are costal inundation and erosion. Still, many coastal systems, especially beach systems, have recovery mechanisms. Resilience levels have a very important role in the beach morphodynamic status and exposure to potential damaging event assessments. The Portuguese West coast has a high wave energetic environment during winter, where waves with 5 year recurrence period can reach 9.2m and storms are frequent. This research aims to access beach hazard and susceptibility to inundation and erosion. Three beach systems were selected and monitored applying sequential profiling methodology over a four year period (2004-2007). Sta. Rita, Azul and Foz do Lizandro beaches are representative systems of the coastal stretch between Peniche and Cascais, which is a cliff dominated coast. Results from the monitoring campaigns are presented, including volume budgets, beach face slope changes, berm occurrence and heights and planimetric coastline dynamics. A hazard and susceptibility assessment diagram and zonation are proposed, including the parameterization of local flood (i.e. mean sea level, maximum spring tides, and surge and run-up levels) and erosion potentials (i.e. volume budget and beach planimetric dynamics).

Keywords: Beach, susceptibility, inundation, erosion, index, extreme events.

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Resumo – Suscetibilidade à inundação e à erosão aplicados a praias expos-TAS. A dinâmica de praia é predominantemente influenciada pelas forças hidrodinâmicas que atuam sobre os sedimentos e determinam quer o seu balanço volumétrico quer a sua dinâmica planimétrica. Estas forças definem a magnitude e a frequência das mudanças morfológicas registadas que, em boa parte do tempo, não implicam qualquer tipo de risco para o sistema ou para as populações. No entanto, as ondas associadas a temporais que incidem sobre as praias são também responsáveis por recuos extremos nos perfis de praia. Estes eventos originam perdas sedimentares muito significativas e o rápido recuo da linha de costa, colocando em risco infraestruturas, propriedades e pessoas. Este risco é frequentemente materializado através da exposição à inundação e à erosão de origem marinha. O litoral ocidental português tem elevada energia hidrodinâmica com ondas com 5 anos de período de retorno cuja altura significativa pode atingir 9,2m. O objetivo deste trabalho é avaliar a exposição de praias à inundacão e à erosão, resultando da ocorrência de eventos hidrodinâmicos extremos. Apresentam-se os resultados da cartografía associada aos índices de potencial de inundação e erosão propostos para 3 sistemas de praia localizados no litoral entre Peniche e Cascais. Para o efeito foram estudadas e monitorizadas as de Sta. Rita, Azul e Foz do Lizandro com recurso à metodologia dos perfis de praia sequenciais durante 4 anos (2004-2007). Os resultados apresentados e a posterior modelação dos perfis em condições extremas de ondulação permitem apresentar um zonamento da suscetibilidade destes sistemas à inundação e à erosão.

Palavras-chave: Praia, suscetibilidade, inundação, erosão, índices, eventos extremos.

Résumé – Susceptibilite aux inondations et a l'erosion des plages exposees. La dynamique des plages est fortement influencée par les forces hydrodynamiques agissant sur les sédiments, ce qui détermine leur bilan volumétrique, ainsi que leur dynamique planimétrique. Ces forces sont responsables de l'ampleur et de la fréquence des modifications morphologiques enregistrées par les plages. Celles-ci, la plupart du temps, ne constituent aucun danger pour le système littoral, ni pour les habitants. Cependant, les houles associées aux tempêtes affectant le littoral, et surtout les plages, sont souvent responsables de retraits accentués des profils de plage, de pertes sédimentaires considérables et d'un recul rapide de la ligne de côte, mettant en danger infrastructures, biens et individus. Ce risque provoque souvent des inondations et des formes d'érosion d'origine marine. Le littoral occidental portugais connait un fort hydrodynamisme, avec des houles de tempête qui atteignent jusqu'à 9,2 m de hauteur significative, avec une période de récurrence de 5 ans. L'objectif de cette étude est d'évaluer l'exposition des plages à l'inondation et à l'érosion qui résultent des événements hydrodynamiques extrêmes. On présente les résultats cartographiques basés sur les indices de potentiel d'inondation et d'érosion, proposés pour 3 systèmes de plages, situés entre Cascais et Peniche. Les plages de S^{1a}. Rita, Azul et Foz do Lizandro ont été étudiées et monitorisées, en utilisant la méthode de profils de plage séquentiels pendant une période de 4 ans (2004-2007). La modélisation ultérieure des profils en situation d'ondulation extrême a permis de cartographier le zonage de susceptibilité aux inondations et à l'érosion de ces systèmes littoraux.

Mots-clés: Plages, susceptibilité, inondation, érosion, indices, événements extrêmes.

I. INTRODUCTION

Beach systems are highly dynamic coastal systems. The main reasons that explain this behaviour are related to the non-cohesive nature of the sediments and their dependence on several triggering factors that have a high energetic potential.

Waves are the main triggering factor acting on the beach profile that produce morphological change through sediment transport.

This means that the main morphological changes associated with sudden sediment loss are, in most cases, attributed to the different types of storm events according to modal wave energy and direction in a short period of time.

Beach hazard assessment is an essential step for susceptibility analysis in these systems. Hazard is understood here as the changes in coastal processes on the beach profile in a frequency/magnitude time scale. However, it must be stressed out that, unlike cliffs, beaches are resilient systems, and their dynamic equilibrium is dependent on the balance between local vulnerability and their recovering capacity.

The most important consequences of extreme wave action over the beach profile are inundation of coastal areas by backshore overwash, erosion by sediment loss and coastline recession causing significant damage to natural systems and property.

The aim of this study is to access beach susceptibility to inundation and erosion on the west coast of Portugal, between *Peniche* and *Cascais* (fig. 1). Beach susceptibility is understood as the spatial incidence of different types of hazard.

II. STUDY SITES

The western coast of Portugal has a semidiurnal mesotidal regime with maximum amplitude above chart datum of circa 4m. The coastal stretch between *Peniche* and *Cascais* is a wave dominated rocky coast, highly conditioned by the Atlantic atmospheric circulation. This results in a clear seasonal differentiation of wave patterns.

The mean annual wave height values registered on the west coast of Portugal vary between 1.8m and 2.4m (Carvalho and Baceló, 1966; Oliveira Pires and Pessanha, 1986; Costa, 1987; Teixeira, 1990; Soares, 1999) and have a clear seasonal differentiation. Winter wave heights reach 2.5m (Oliveira Pires and Pessanha, 1986) when the total west coast is considered. Vitorino *et al.* (2002), also stated that the mean summer wave significant height in the northwestern coast is between 1m and 2m; however, when the mean winter values are analysed this parameter reaches 3.5m. The dominant wave direction is NW (265 days/year, Costa, 1994).

A storm is defined to occur on the Portuguese West coast when a 5m offshore significant wave height (H_s) is registered (Pita and Santos, 1989; Costa, 1994; Gama *et al.*, 1994).

The *Peniche – Cascais* coastal sector (fig. 1) is lacking in sediment supply from longshore drift and local sources because of: (i) the dominantly limestone nature of the cliffs; (ii) the presence of the *Nazaré* Canyon and the *Peniche* headland, two natural obstacles to sediment transport by longshore drift; (iii) the small sized river basins that are unable to effectively contribute to regional sediment budget.

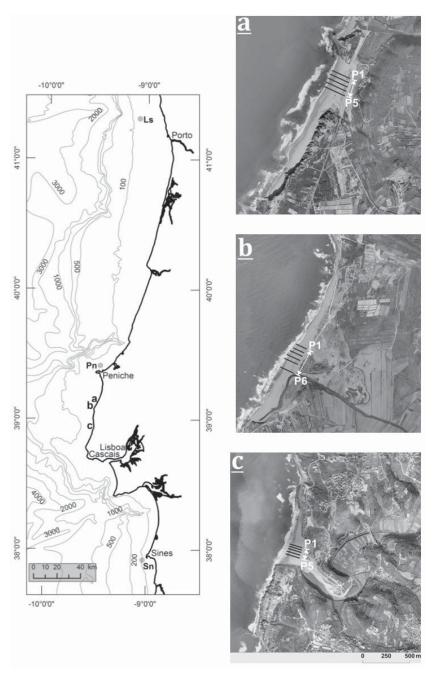


Fig. 1 – Study area. S^{ta}. Rita beach (a), Azul beach (b), Foz do Lizandro beach (c) and wave buoys at Leixões (Ls), Peniche (Pn) and Sines (Sn). (See coloured version online)

Fig. 1 – Área de estudo. Praia de S^{ta}. Rita (a), praia Azul (b), praia da Foz do Lizandro (c) e bóias ondógrafo em Leixões (Ls), em Peniche (Pn), e em Sines (Sn).

In this coastal stretch beach systems are narrow, embayed and/or part of small estuaries.

The study sites of *S*^{*ia*}. *Rita*, *Azul* and *Foz do Lizandro* beaches (fig. 1) are equally exposed W to NW waves but distinct in size, shape and type (table I).

 S^{ta} . Rita beach is a composite beach system with two components of beach-dune in its northernmost part and a beach-cliff system at the southern end. The dune field sub-component shows signs of heavy anthropogenic degradation, with a dense trampling network and abundant deflation surfaces. Cliff sub-component dynamics are mainly dependent on continental erosion processes (table I).

	Table I – Type and main beach system characteristics.
Quadro	I – Definição do tipo e características dos sistemas de praia em análise.

Beach	S ^{ta} . Rita	Azul	Foz do Lizandro
Present day system type	Beach – Dune / Beach – cliff	Beach – Dune	Beach – Dune
Pre-existing system type	Spit / Coastal lagoon	Wide alluvial plain / Ria	Narrow estuary
Slope of the internal continental shelf $(0m \grave{a} \rightarrow -40m)$	0.33°	0.36°	0.61°
Subaerial beach area	113 019.27m ²	140 304.22m ²	93 126.07m ²
Total subaereal system area	150 760.80m ²	569 785.31m ²	120 605.68m ²

Azul beach-dune system is the largest system among the three study sites reaching 1900m along the coastline. Beach sub-component is widely exposed to NW dominant wave climate and to the W and SW storm waves, resulting in a modal steep slope profile. The dune field sub-component is the second largest in the *Peniche-Cascais* coastal stretch and occupies a wide alluvial plain (table I).

Foz do Lizandro beach is the smallest one and is a beach-dune system type with 600m of coastal length. Beach and dune components are highly anthropogenic intervened (table I).

III. METHODOLOGY

1. Susceptibility analysis conceptualization

Inundation and erosion of coastal areas are among the most frequent problems that stakeholders have to deal with and, generally, a spatial definition of exposed areas to hazard (susceptibility analysis) is needed in the integrated coastal zone management process.

Susceptibility analysis is calculated and mapped based on the inundation potential index and on the erosion potential index (fig. 2) (Trindade, 2010). The main conditioning and triggering factors are taken into account, in particular those related with marine hydrodynamics, system morphodynamics and sediment characteristics.

1.1. Inundation potential index

Beach and backshore flooding are triggered in the first place by overwash events that can last from minutes or hours to days, according to the storm energetic content and period. Overwash occurrence is dependent on incident wave characteristics, tide height, storm surge height, runup levels and, ultimately and of most importance, beach slope and berm presence (fig. 2).

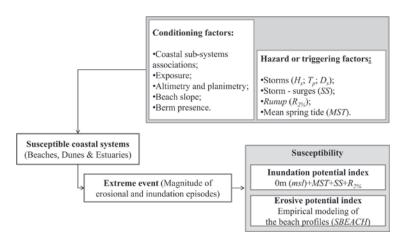


Fig. 2 – Wave dominated beach susceptibility conceptualization.

Fig. 2 – Conceptualização da avaliação da susceptibilidade em praias expostas.

Inundation potential (*IPI*) was calculated for the three beach systems (eq. 1) by successively summing the different water levels in near reflective $(tan\beta_{max})$ and dissipative $(tan\beta_{min})$ profiles morphology. Water levels considered are the following:

- (i) Mean sea level (0m msl) which is the reference base level obtained in the Cascais tide gauge;
- (ii) Mean spring tide (MST) obtained from one year measurements in the Cascais tide gauge;
- (iii) Storm surge maximum levels (SS_{max}) calculated from the differences between the high accurate records of water levels (3 minutes frequency) from the Portuguese Geographical Institute in the *Cascais* tide gauge and the predicted tides for the same area from the Hydrographic Institute, over the reference period (January 2004 July 2007);
- (iv) Runup values ($R_{2\%}$) are calculated with Stockdon *et al.* (2006), formulations for intermediate to reflective profiles (eq. 2; $\xi_0 > 3$) and dissipative profiles (eq. 3; $\xi_0 < 3$), in every beach profile campaigns. For ξ_0 values and beach morphodynamic classification the eq. 4 is applied.

$$IPI_{tan\beta \ min}^{tan\beta \ max} = 0m \ (msl) + MST + SS_{max} + R_{2\%}$$
 Eq. 1

$$R_{2\%} = 1.0 \left(0.35 \beta_f (H_0 L_0)^{0.5} + \frac{\left[H_0 L_0 (0.563 \beta_f^2 + 0.004) \right]^{0.5}}{2} \right)$$
 Eq. 2

$$R_{2\%} = 0.0043(H_0L_0)^{0.5}$$
 Eq. 3

$$\xi_0 = \frac{\beta_f}{\frac{H_0}{L_0}^{0.5}}$$
 Eq. 4

were H_0 is the offshore mean significant wave height of the weekly period prior to each campaign, L_0 is the correspondent wave length, β_f is the mean beach face and low tide terrace slope, because these are the most exposed morphological elements of the beach profile to waves and currents.

1.2. Erosion potential

Beach erosion phenomena can be attributed to natural or anthropogenic causes, acting in very short time scales (storms or tsunamis) or more long-term trends (sea level rise). One of the most important triggering factors for fast sediment loss in the beach systems are extreme events such as storms, because of the associated high frequency and magnitude. Coastline recession, foredune erosion and beach and dune crest lowering are among the most important consequences of these episodes.

Highly accurate and reliable measurements of the beach response to these extreme energetic events are crucial for mean behaviour characterisation and extreme erosional response prediction, through numerical modeling of the profile morphological modifications.

In this study we use SBEACH numerical model (Larson and Kraus, 1989; Larson et~al., 1990; Rosati et~al., 1993; Wise et~al., 1996 and Larson and Kraus, 1998) for extreme beach profile behaviour prediction, allowing assess to coastline recession over a 100 year return period storm. SBEACH is an empirical based numerical model which allows the use of calibration parameters (K~and~C) for best fitting model results to local conditions of sediment transport and incident wave field. Model application follows in a two moment basis, with a first calibration approach and a second validation process. This model is also frequently used to predict erosion over the beach profile under extreme wave occurrence (Farrar et~al., 1994; Schoonees and Theron, 1995; Zheng and Dean, 1997; Carley et~al., 1999;

Fauver, 2005; Tinh, 2006; Fauver et al., 2007; King Jr., 2007; Canizares and Irish, 2008; Creed and Reilly, 2008; Ponce, 2008; Corbau et al., 2009; Jimenez et al., 2009).

Before and after storm profiles are required for model calculations, as well as wave and tide data.

Two storm events per beach were chosen for model calibration and validation, allowing to access SBEACH performance both spatially, with the validation results for each profile in each beach, and in time. A total of 10 profiles in S^{ta} . Rita, 12 in Azul and 10 in Foz do Lizandro, where used in this process.

The calibration process was applied only to the first profile in each beach (P1 - fig. 1) and to the first storm event, using morphodynamic and hydrodynamic reference terms to locally calibrate the empirical model parameters K (transport rate) and C (slope dependent term).

The validation process consisted in the systematic application of calibrated K and E values to predict volume loss from the remains profiles (P2 to P5 or P6 – fig. 1) in each beach and the comparison of predicted volumes with measured ones.

2. Data acquisition

A systematic beach profile survey was carried out over a four year period (2004-2007), resulting in 279 topographic transepts distributed over the three beach systems including emerged beaches and foredune subsystems. The total number of profiles ready for survey in each beach system per campaign is: *S^{ta}*. *Rita* beach 5 profiles; *Azul* beach 6 profiles, *Foz do Lizandro* beach 5 profiles (fig. 1). Berm and beach face dynamics, including crescentic beach morphology, determined the profiles spacing and position.

dGPS and total station were used in fixed beach profile monitoring and an overlap of the transepts over time was assured. The differences between measurements acquired by both equipments proved to be negligible in this dynamic context (Trindade *et al.*, 2007).

Between September 2005 and July 2006, the surveys were preformed bimonthly, allowing the detailed record of several pre- and post-storm morphology data. This is essential for the assessment and modeling of the system response to high energetic events in terms of volumetric budget and planimetric dynamics. Measured volumetric budget was calculated above mean sea level and bellow the point of no relative hydrodynamic sand movement, excluding dune morphology changes due to aeolian transport.

Superficial sediment sampling was performed as part of each campaign data collection in order to evaluate the characteristic grain size of each system. Samples were collected from the most dynamic morphological elements of the beach

profile, namely the berm, the beach face, the base of the beach face and the low tide terrace. Approximately 60g of sediment was washed and dry-sieved; sand statistics parameters were calculated according to the logarithmic method of the moments with *Gradistat* spreadsheet (Blott and Pye, 2001). Sediment size is a very important parameter in beach profile modeling because it defines beach slope and infiltration/exfiltration processes in the beach profile.

Wave parameterization based on significant heights (H_s) , periods (T_s) and directions (D_s) was only available from the two west coast wave buoys still functioning, the Leixões and Sines buoys (fig. 1). This data was used to assess storm frequency and magnitude during the monitoring period, establishing the basic hydrodynamic parameterization for profile model calibration and validation. Significant wave return periods were calculated from the nearest *Peniche* wave buoy 1980-1985 time series (fig. 1), using generalised extreme value distribution with R^{\odot} open source software.

Tide data was obtained from public secular data series of water heights at the *Cascais* tide gauge. Storm surge values were evaluated using the difference between the Portuguese Geographical Institute measured *Cascais* tide gauge records and the Hydrographic Institute tide height predictions for every storm event registered in the *Leixões* and *Sines* wave buoys during the reference period (January 2004 to July 2007, tables II and III). Both data series are referenced to the *Cascais* Datum which means that sea level change in the last decades (≈1.9mm yr-1, Antunes and Taborda, 2009) is not considered in this study. Consequently, modeled water level calculations may express a slight underestimation.

IV. RESULTS

Hydrodynamic, morphodynamic and sediment data is now presented following the conceptual model described above.

1. Inundation potential index

1.1. Storm surge levels

It was found that both the values above the 90^{th} percentile and the maximum storm surge values (tables II and III) were very high when compared to other reference studies in this area (Taborda and Dias, 1992 - 0.50m; Gama *et al.*, 1994 - 0.52m).

Table II – Storm surge values of the storm events registered in the Leixões wave buoy. $(\geqq P_{k90} \text{ values underlined}) \\ H_s - \text{ significant wave height; } H_{\textit{max}} - \text{maximum wave height; } D_s - \text{ significant wave direction.}$

Quadro II – Valores de sobre-elevação do nível do mar de origem meteorológica dos temporais registados na bóia ondógrafo de Leixões (valores $\geqq P_{k90}$ sublinhados); H_s – altura significativa das ondas; H_{max} – altura máxima das ondas; D_s – direcção significativa das ondas.

Storm beginning day	Storm surge (m)	Storm tide cycles	H_s (m)	H_{max} (m)	$D_s $ (°)
08.01.2004	0.19	6	4.38	8.91	292.5 (90.6%)
02.04.2004	0.43	2	4.85	10.12	315.0 (100.0%)
08.10.2004	0.44	1	4.44	9.43	225.0 (36.4%)
27.10.2004	0.65	4	4.25	10.09	315.0 (68.8%)
19.12.2004	0.26	2	4.63	9.10	315.0 (100.0%)
25.12.2004	0.25	3	5.26	10.69	315.0 (100.0%)
18.01.2005	0.25	3	5.92	13.50	315.0 (98.6%)
25.03.2005	0.27	8	4.60	9.98	270.0 (90.0%)
02.12.2005	0.37	3	4.86	12.02	315.0 (74.3%)
01.01.2006	0.40	2	4.35	9.77	315.0 (92.7%)
06.01.2006	0.40	1	4.79	9.86	315.0 (83.3%)
15.01.2006	0.22	2	4.90	9.81	292.5 (86.7%)
21.05.2006	0.35	10	5.59	10.77	292.5 (73.3%)
23.10.2006	0.78	3	5.54	12.01	292.5 (100.0%)
24.11.2006	0.55	2	4.75	11.54	292.5 (40.8%)
03.12.2006	0.34	3	4.61	10.62	315.0 (80.3%)
07.12.2006	0.34	1	6.01	13.34	315.0 (73.8%)
09.02.2007	0.27	1	4.83	12.61	292.5 (62.2%)
22.02.2007	0.39	1	5.22	11.70	292.5 (93.8%)
06.03.2007	0.14	2	4.67	10.76	315.0 (51.1%)
19.03.2007	0.16	1	4.95	10.30	337.5 (52.1%)
Maximum	0.78				

Table III – Storm surge values of the storm events registered in the Sines wave buoy. ($\geqq P_{k90}$ values are underlined)

Quadro III – Valores de sobrelevação do nível do mar de origem meteorológica dos temporais registados na bóia ondógrafo de Sines.

(valores ≧P_{to0} encontram-se sublinhados)

Begining day	Storm surge (m)	Tide cycles	<i>H</i> _s (m)	H _{max} (m)	D _s (°)
25.03.2005	0.27	8	4.78	9.77	292.5 (52.8%)
18.02.2006	0.27	1	4.80	9.98	315.0 (60.5%)
24.11.2006	0.55	3	4.57	11.84	292.5 (39.6%)
05.12.2006	0.34	1	4.44	9.62	315.0 (80.6%)
08.12.2006	0.34	1	4.76	10.58	315.0 (65.5%)
10.02.2007	0.27	1	4.38	9.77	315.0 (78.2%)
22.05.2007	0.39	1	5.04	11.58	292.5 (98.3%)
Maximum	0.55				

Storm surge data analysis revealed that these events lasted no longer than 10 tide cycles (approximately 5 days). Considering the most extreme events ($P_{k90} = 0.55 \text{m}$), storm surge can be considered a very brief episode lasting no longer than 4 tide cycles.

No correlation was found between storm waves and storm surge magnitudes.

1.2. Runup levels

Runup levels (Eq. 2 and 3) were calculated as a function of the maximum observed wave data during the reference period (maximum reference level, table IV) and based on a 100 year return period storm data (maximum predicted level, table IV) for the lower and higher slope values registered during the monitoring campaigns (table IV). As expected the higher runup levels are associated to near reflective behaviour of the beach profile. In these type of profiles energy dissipation occurs near the coastline leading to higher levels of setup and runup. The discriminated calculation of levels according to the morphodynamic state of the beach is, therefore, essential for susceptibility analysis.

Table IV – Maximum flood levels for near reflective and dissipative profiles in S^{ta} . Rita (SR), Azul (AZ) and Foz do Lizandro (LZ) beaches.

 $R_{2\%}$ maximum reference level – maximum observed H_s storm;

 $R_{2\%}$ maximum predicted level – extreme storm with a 100 year return period; msl – mean sea level; L_{θ} – offshore wave length;

 SS_{max} – maximum storm surge level; MST – mean spring tide.

Quadro IV – Níveis máximos de inundação associados a perfis próximos do refletivo e dissipativa, para as praias de S^a. Rita (SR), Azul (AZ) e Foz do Lizandro(LZ).

 R_{296} maximum reference level -valor máximo de H_{2} observada;

 $R_{2\%}$ maximum predicted level – valor previsto de H_s para um temporal com per. retorno de 100 anos; msl – nível médio do mar; L_o – comprimento de onda de largo;

SS_{max} – Nível máximo de storm surge observado; MST – nível médio de preia-mar.

		$taneta_{_{min}}$			$taneta_{_{max}}$			
		SR (0.03)	AZ (0.04)	LZ (0.03)	SR (0.08)	AZ (0.09)	LZ (0.09)	
R _{2%}	Maximum reference level $(H_s = 8.70 \text{m}; L_0 = 206.93 \text{m})$	2.06	2.29	2.06	3.34	3.63	3.63	
(m)	Maximum predicted level $(H_s = 12.38 \text{m}; L_0 = 506.07 \text{m})$	3.84	4.27	3.84	6.63	6.77	6.77	
	SS _{max} (m)			0.	78			
	MST (m - msl)			1.:	59			

2. Erosion potential index

2.1. Morphodynamic and hydrodynamic reference values for model calibration and validation

Two episodes per beach system of measured significant sediment loss in a short period (typically 15 days) were selected as a morphodynamic reference basis for model testing. Volume budgets per profile are almost always negative and this is more evident in the *Foz do Lizandro* system (table V).

Hydrodynamic reference values per beach show very distinct episodes in the duration of events and significant wave height distribution over time (fig. 3).

Reference storms for S^{ta} . Rita beach were the most severe, with H_s reaching 7.34m.

Table V – Morphodynamic reference values related to two storm occurrence (a and b).

Quadro V – Valores morfodinâmicos de referência relacionados com a ocorrência de dois temporais (a e b).

Beach		Days		P	rofile volume	budget (m³/n	n)	
			P1	P2	Р3	P4	P5	Р6
CD	а	15.02.2006 ↓ 01.03.2006	-264.5	-167.7	-202.8	-158.3	141.5	
SR	b	14.06.2006 ↓ 28.11.2006	-278.7	-162.5	-204.7	-9.2	-50.0	
47	а	31.01.2006 ↓ 14.02.2006	-190.1	-220.8	-252	-187.3	-168.5	-477.1
AZ	b	28.02.2006 ↓ 14.03.2006	-67.3	-225.8	-117.1	-72.4	-38.2	-236.2
17	а	13.02.2006 ↓ 02.03.2006	-106.1	-204.5	-135.7	-524.5	-521.8	
LZ	b	02.03.2006 ↓ 15.03.2006	-735.9	-706.8	-843.1	-476	-347.5	

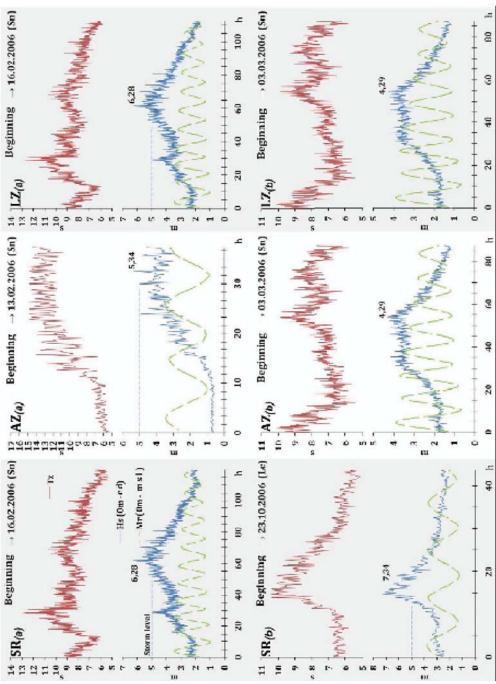


Fig. 3 – Hydrodynamic reference values related to two storm occurrence (a and b).

Le – Leixões wave buoy; Sn – Sines wave buoy; cd – chart datum; msl – mean sea level; Tz – peak period;

Hs – significant wave height; Mr – tide height. (See coloured version online)

Fig. 3 – Valores hidrodinâmicos de referência relacionados com a ocorrência de dois temporais (a e b).

Le – bóia ondógrafo de Leixões; Sn – bóia ondógrafo de Sines; cd – zero hidrográfico; msl – nível médio do mar; Tz – período de pico das ondas; Hs – altura significativa das ondas; Mr – altura da maré.

2.2. Model calibration

The calibration process involved the successive adjustment of the K (transport rate term) and \mathcal{E} (slope dependent term) for P1 in each beach system and on the first storm event to best fit measured results. Default values of K and \mathcal{E} were then replaced by better fitted ones and results are expressed in table VI. In the worst case scenarios, which are the results of S^{ta} . Rita beach, the model explains over 89.1% of the measured profile, leaving 10.9% of unexplained volume change data (table VI).

Table VI – Local calibration values for P1 related to measured profiles and predicted data from SBEACH model.

dro VI – Valores de calibração local para P1 e comparação entre volumes medidos e previs

Quadro VI – Valores de calibração local para P1 e comparação entre volumes medidos e previstos
pelo modelo SBEACH.

	SR (P1)	AZ (P1)	LZ (P1)
Storm (beginning)	16.02.2006	13.02.2006	16.02.2006
Measured pre-storm volume (m³/m)	173.87 (15.02.2006)	90.59 (31.01.2006)	340.12 (13.01.2006
Measured post-storm volume (1) (m³/m)	147.42 (01.03.2006)	71.58 (14.02.2006)	329.51 (02.03.2006)
Measured volume budget (2) (m³/m)	-26.45	-19.01	-10.61
K (adjusted) (m ⁴ /N)	2.50*10-7	2.50*10-7	2.50*10-7
€ (adjusted) (m²/s)	0.005	0.005	0.005
Predicted post-storm volume (m³/m)	131.41	69.90	301.64
Predicted volume budget (3) (m³/m)	-42.46	-20.68	-38.48
Relative budget difference (((3-2)*100)/1) (%)	-10.9	-2.3	-8.5
Budget difference 3-2 (m³/m)	16.01	1.68	27.87

2.3. Model validation

K and C calibrated values were then systematically applied to the rest of the 29 profiles in the three beach systems to access model performance and results validation. Results are shown in figure 4 and table VII.

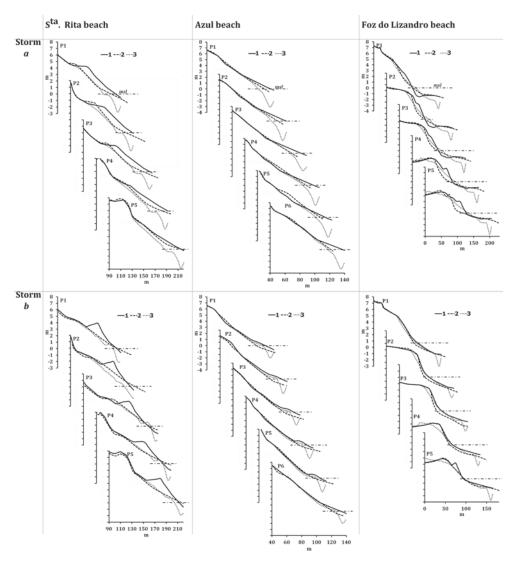


Fig. 4 – Beach profiles related to storm occurrence (*a* and *b*). Pre-storm measured (1), post-storm measured (2) and predicted/modelled (3) *msl* – mean sea level.

Fig. 4 – Perfis de praia pré-temporal medidos (a), pós-temporal medidos (b) e previstos/modelados (c) relacionados com a ocorrência de temporais (a e b). msl – nível médio do mar.

The profile modeling results (fig. 4) show good corresponding results above *msl* considering the measured post-storm and the predicted profiles. Poor correspondence below *msl* is related to lack of bar morphodynamic data and for this reason volume change was systematically calculated above *msl*. The *Foz do Lizandro* beach shows a slight erosion overestimation on berm morphology. The same slight erosion overestimation occurs in *S^{ta}*. *Rita* beach on P5 (storm a) and on P2 (storm b).

The relative volume difference between post-measured and predicted volumes can be assumed as a significant level parameter. In general we can assume the good performance of the model (table VII – relative volume difference). In most cases (20 profiles in 32) the relative volume difference does not exceed 10%, which means that the model is explaining in most cases more than 90% of the measured changes due to storm events. The maximum relative difference value is -39.7%.

Table VII –	Validation results and applicability of the <i>SBEACH</i> model.
Quadro	VII – Validação e aplicabilidade do modelo SBEACH.

Storm	Profile	Measured post-storm volume (m³/m)			Pr	edicted pos volume (n			Relative vo difference	
		SR	AZ	LZ	SR	AZ	LZ	SR	AZ	LZ
	P1	147.4	71.6	329.5	131.4	69.9	301.6	-10.9	-2.3	-8.5
	P2	223.6	76.5	298.9	179.3	72.5	282.7	-19.8	-5.2	-5.4
	Р3	157.5	79.5	281.3	127.7	85.5	260.8	-18.9	7.5	-7.3
a	P4	187.7	87.7	221.4	150.3	86.2	220.5	-19.9	-1.6	<u>-0.4</u>
	P5	158.2	95.8	206.7	95.4	91.7	201.7	<u>-39.7</u>	-4.3	-2.4
	P6	-	116.6	-	-	127.1	-	-	9.0	-
•••••	P1	169.3	63.4	255.9	149.8	59.9	238.2	-11.5	-5.5	-6.9
	P2	247.3	63.0	228.2	199.8	68.0	215.8	-19.2	8.0	-5.4
	Р3	184.7	68.1	197.0	169.6	64.6	201.5	<u>-8.2</u>	-5.2	2.3
b	P4	207.3	75.3	173.8	183.9	66.6	157.1	-11.3	-11.5	-9.6
	P5	143.0	83.5	172.0	126.0	71.3	142.9	-11.9	<u>-14.7</u>	<u>-16.9</u>
	P6	-	122.9	-	-	123.6	-	-	0.6	-

V. BEACH SUSCEPTIBILITY MAPPING FOR WAVE DOMINATED BEACHES

Mapping beach susceptibility to erosion and flooding based on field data measurements and profile modeling revealed different results in the three beach systems. Results of predicted flooding areas and beach recession are analysed within the system limits.

1. Sta. Rita beach

Coastline recession in the two proposed scenarios (initial dissipative or reflective profiles) have slight consequences on the beach dune component of the system (fig. 5a). Even the 3m contour recession, which may be assumed as the berm crest or the boundary between the foreshore and the backshore, have no significant consequences in the dune sub-system, considering a 100 year storm return period. Unlike this system component, the southernmost beach-cliff sector may experience severe erosion caused by beach vanishing. Continental predominant geomorphologic processes may be replaced by marine erosion. Predicted flooding levels also have no important impact neither inside nor outside the system boundaries (fig. 5b). Contrasted flooding areas between the north and the central sectors of this beach system are due to the presence and morphology of the dune field in the northern sector. The central sector of the *Sta. Rita* beach system is occupied by a very small and narrow alluvial plain (*Sorraia* brook), which conditions the inland extent of the predicted flood area.

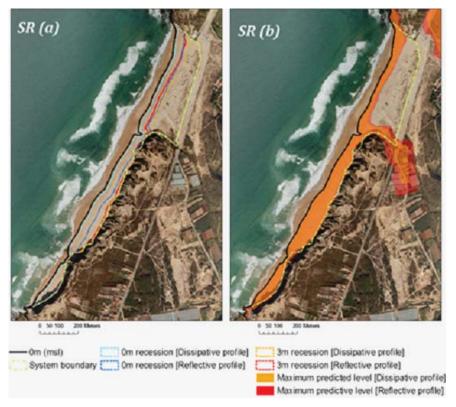


Fig. 5 – Erosion [SR(a)] and inundation [SR(b)] potential indexes zonation for S^{ta} . Rita beach. (See coloured version online)

Fig. 5 – Cartografia do potencial de erosão [SR(a)] e de inundação [SR(b)] na praia de S^{ta} . Rita.

2. Azul beach

Predicted erosion and coastline recession will cause the disappearance of the emerged beach, with extreme foredune erosion (fig. 6a). This is particularly problematic because the southernmost spit sector will probably be replaced by an open estuary endangering the *Foz* locality. Extreme inundation levels confirm the high risk exposure of this locality to generalized overwash processes and direct wave damage (fig. 6b).

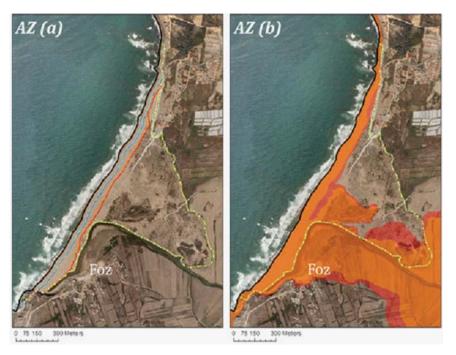


Fig. 6 – Erosion [AZ(a)] and inundation [AZ(b)] potential indexes zonation for Azul beach (legend on fig. 5). (See coloured version online)

Fig. 6 – Cartografia do potencial de erosão [AZ(a)] e de inundação [AZ(b)] na praia Azul. (legenda na fig. 5)

3. Foz do Lizandro beach

Severe berm level recession will replace the present closed estuary system with an open one with consequences in the actual beach support structures (fig. 7a). The beach sub-system area will be drastically reduced, but the small dune field is well protected from erosion due to its sheltered position. Extreme flood levels associated with marine hydrodynamics will affect the total area of the estuary floodplain (fig. 7b), with high potential of risk to the local population and properties.

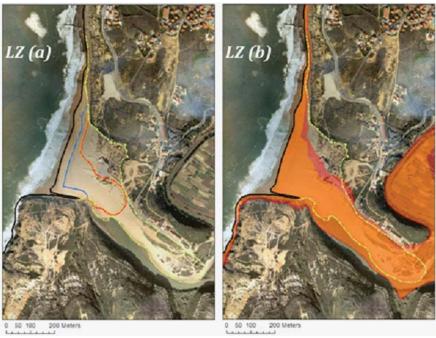


Fig. 7 – Erosion [LZ(a)] and inundation [LZ(b)] potential indexes zonation for *Foz do Lizandro* beach (legend on fig. 5). (See coloured version online)

Fig. 7 – Cartografia do potencial de erosão [AZ(a)] e de inundação [AZ(b)] na praia da Foz do Lizandro. (legenda na fig. 4).

VI. CONCLUDING REMARKS

Like other systems in wave dominated costal environments, beaches are exposed to extreme hydrodynamic events that can cause sudden remarkable modifications in the morphology, often causing a high magnitude of sediment loss and coastline retreat. Predicting this extreme system behaviour, through erosion or inundation processes, can prevent infrastructure and property damage and, in limited situations, loss of life.

Susceptibility mapping methodology of the predicted areas of extreme erosion (coastline retreat) and inundation proved to be applicable to the studied beach systems.

The inundation potential index show that a series of marine water levels can be combined based on the existing time series of waves to achieve a 100yr return period levels. Because runup levels are slope dependent, it is of fundamental interest to map the inundation parameters associated with dissipative and reflective profiles. It was shown that results can be considerable different according to the initial type of beach profile.

The erosion potential index revealed the importance of the beach extreme profile behaviour modeling and empirical adjustment. The application of the *SBEACH* model proved to be a very versatile tool for extreme beach retreat prediction. The calibration process and especially the validation results show good adjustment between predicted and measured values of beach retreat under storm conditions.

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BIBLIOGRAPHY

- Antunes C, Taborda R (2009) Sea level at Cascais tide gauge: data, analysis and results. *Journal of Coastal Research*, SI 56 (Proceedings of the 10th International Coastal Symposium). Lisbon: 218-222.
- Blott S J, Pye K (2001) Gradistat: a grain size distribution and statistics package for the analysis of unconsolidated sediments. Earth Surf. Process. Landforms 26: 1237-1248.
- Canizares R, Irish J (2008) Simulation of storm-induced barrier island morphodynamics and flooding. *Coastal Engineering*, 55: 1089-1101.
- Carley J, Turner I, Couriel E, Jackson L, Mcgrath J (1999) The practical application of four commercially available numerical beach morphology models on a high energy coastline. Proceedings of Coasts and Ports '99, Perth, Australia. Institute of Engineers, Australia: 101-106.
- Carvalho J, Barceló J (1966) Agitação marítima na costa W de Portugal. Memórias do L.N.E.C., 290.
- Corbau C, Simeoni U, Archetti R, Peretti A, Farina, M (2009) Winter sandy protections of the Northern Adriatic coast against flooding: preliminary results. *Journal of Coastal Resear*ch, SI56: 1194-1198.
- Costa M (1994) Agitação marítima na costa portuguesa. Anais do Instituto Hidrográfico, 13: 35-40.
- Costa M (1987) Análise de um ano de dados de agitação marítima na Figueira da Foz. *Anais do Instituto Hidrográfico*, 8: 23-28.

- Creed C, Reilly W (2008) Site-Specific Calibration of the SBEACH Model for a Southeast Florida Beach. Proceedings from the 21st National Conference on Beach Preservation Technology. Sarasota, Florida.
- Farrar P, Borgman L, Glover L, Reinhard R, Pope J, Swain A, Ebersole B (1994) Storm impact assessment for beaches at Panama City, Florida. Technical Report CERC-94-11, U.S. Army Corps of Engineers, Vicksburg, U.S.A.
- Fauver L (2005) Toward predicting barrier island vulnerability: simple models for dune erosion. MSc thesis, College of Marine Science, University of South Florida.
- Fauver L, Thompson D, Sallenger A (2007) Modeling dune response using measured and equilibrium bathymetric profiles. *Proceedings of the 6th International Conference on Coastal Sediments*, ASCE, New Orleans: 897-910.
- Gama C, Dias J, Ferreira Ó, Taborda R (1994) Analysis of storm surge in Portugal, between June 1986 and May 1988. Proceedings of Littoral, 94: 26-29.
- Jimenez J, Ciavola P, Balouin Y, Armaroli C, Boscom E, Garvais M (2009) Geomorphic coastal vulnerability to storms in microtidal feetch-limited environments: application to NW Mediterranean & N Adriatic seas. Journal of Coastal Research, SI56: 1641-1645.

- King J R (2007) Wave and beach processes modeling for Sabine pass to Galveston Bay, Texas, Shoreline erosion feasibility study. ERDC/ CHL TR-07-6, Coastal and Hydraulics Laboratory, U.S. Army Engineer Research and Development Center, Vicksburg.
- Larson M, Kraus N (1998) Numerical model for simulating storm induced beach change. Report 5: Representation of nonerodible(hard) Bottoms. Technical Report CERC-89-9, U.S. Army Corps of Engineers, Vicksburg, U.S.A.
- Larson M, Kraus N (1989) Numerical model for simulating storm induced beach change. Report 1: Empirical foundation and model development. Technical Report CERC-89-9, U.S. Army Corps of Engineers, Vicksburg, U.S.A
- Larson M, Kraus N, Byrnes M (1990) Numerical model for simulating storm induced beach change. Report 2: Numerical formulation and model tests. Technical Report CERC-89-9, U.S. Army Corps of Engineers, Vicksburg, U.S.A.
- Oliveira Pires H, Pessanha V (1986) Wave power climate of Portugal. *Proceedings IUTAM Symposium, Hidrodynamics of ocean wave-energy utilization*: 157-167.
- Oliveira Pires H, Pessanha L (1982) Características direccionais da agitação marítima na região de Sines. Instituto Nacional de Meteorologia e Geofísica, Lisboa.
- Pita C, Santos J (1989) Análise dos temporais da costa oeste de Portugal Continental. Report 1/89-A, Instituto Hidrográfico/L.N.E.C., Lisboa.
- Ponce E (2008) Coastal vulnerability to storms in the Catalan coast. PhD thesis, Laboratori d' Enginyeria Maritima, Universitat Politecnica de Catalunya.
- Rosati J, Wise R, Kraus N, Larson M (1993) Numerical model for simulating storm induced beach change. Report 3: User's Manual. Technical Report CERC-93-2, U.S. Army Corps of Engineers, Vicksburg, U.S.A.
- Schoonees J, Theron A (1995) Evaluation of 10 cross-shore sediment transport/morphological models. *Coastal Engineering*, 25: 1-41.

- Soares A (1999) Caracterização do clima de agitação marítima em Portugal continental utilizando os resultados do modelo numérico MAR3G.2. Nota técnica VAM2 – 2/99, Instituto de Meteorologia.
- Stockdon H, Holman R, Howd P, Sallenger J R A (2006) Empirical parameterization of setup, swash, and runp. Coastal Engineering, 53: 573-588.
- Taborda R, Dias J (1992) Análise da sobreelevação do mar de origem meteorológica durante os temporais de Fevereiro/Março de 1978 e Dezembro de 1981. Geonovas, 1: 89-97.
- Teixeira S (1990) Dinâmica das praias da Península de Setúbal (Portugal). Dissertação de Mestrado, Universidade de Lisboa.
- Tinh N (2006) Modelling of coastal overwash. Review and application of state-ofart empirical, numerical and analytical methods. MSct Thesis, Lund Institute of Technology/Lund University.
- Trindade J (2010) Dinâmica morfossedimentar de praias dominadas por sistemas de arriba (Peniche-Cascais). Tese de Doutoramento em Geografia, especialidade de Geografia Física, Universidade Aberta, Lisboa.
- Trindade J, Ramos-Pereira A, Neves M, Borges B, Paixão R (2007) Monitorização da morfodinâmica de praias com dGPS e Estação Total. Comparação de resultados. Publicações da Associação Portuguesa de Geomorfólogos, V: 293-301.
- Vitorino J, Oliveira A, Jouanneau J, Drago T (2002) Winter dynamics on the northern Portuguese shelf. Part 1: physical processes. *Progress in Oceanography*, 52: 129-153.
- Wise R, Smith S, Larson M (1996) SBEACH: Numerical model for simulating storm-induced beach change. Report 4: Cross-shore transport under random waves and model validation with SUPERTANK and field data. Technical Report CERC-89-9, US Army Corps of Engineers, Vicksburg, U.S.A.
- Zheng J, Dean R (1997) Numerical models and intercomparisons of beach profile evolution. Coastal Engineering, 30: 169-201.